Your data and "A priori" Calibration

Rosita Paladino INAF Istituto di Radioastronomia Italian Node of ALMA Regional Center

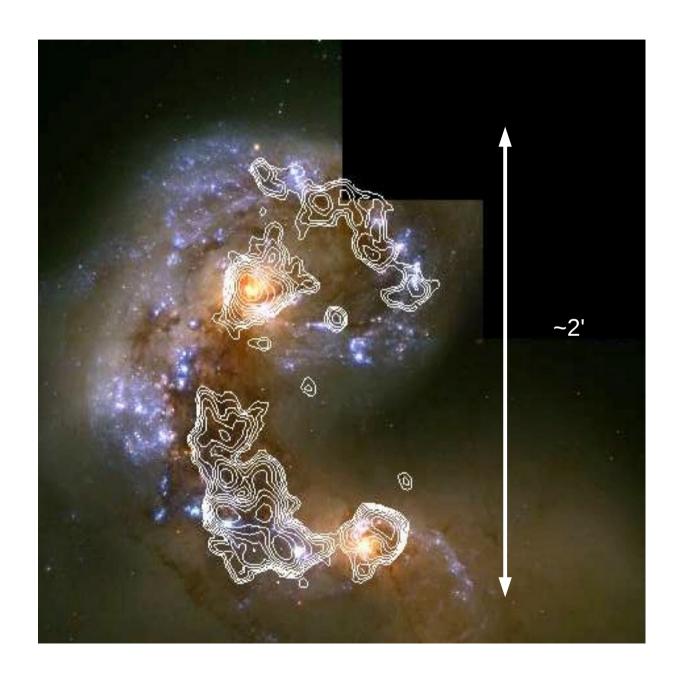
http://www.alma.inaf.it/index.php/Courses

NGC4038/4039



Nearby (z=0.005688) interacting galaxies: NGC4038 & NGC4039

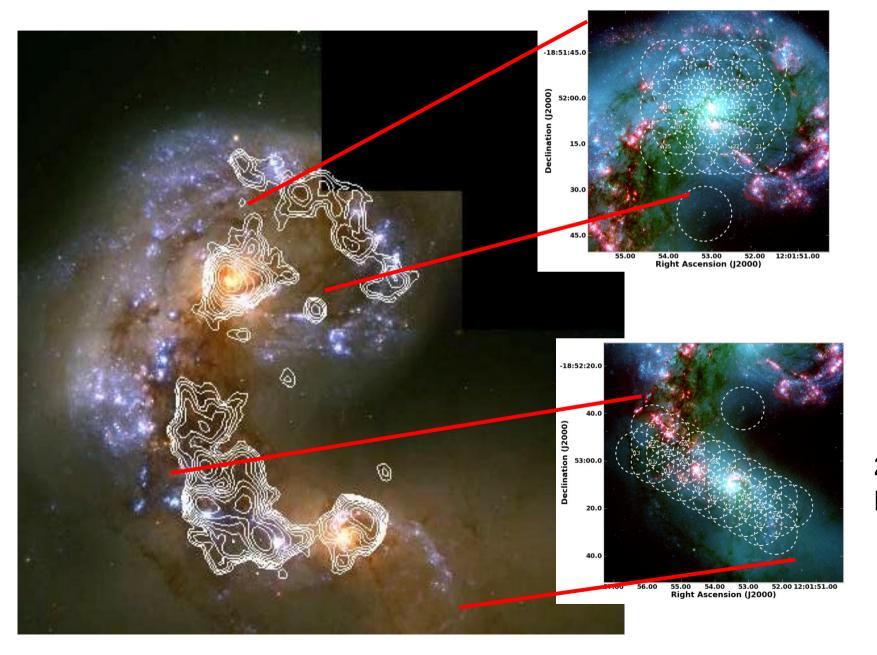
NGC4038/4039



Wilson et al. (2000)

Observations of CO(1-0) resolution 3"x4"

ALMA field of view in Band 7 ~ 15 "



23 pointings

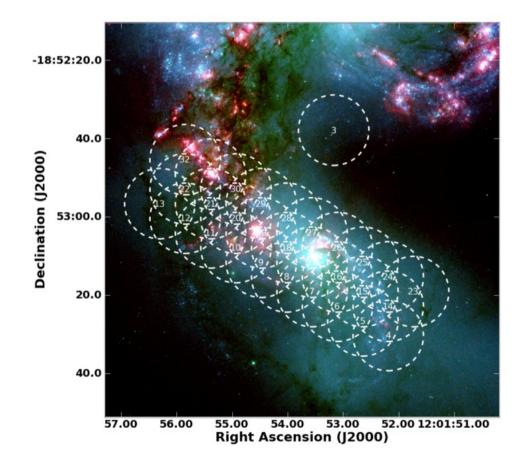
29 pointings

Antennae ALMA SV

ALMA Science Verification data targeting the CO (3-2) line (rest frequency = 345.7960 GHz)

ALMA field of view ~ 15 " ----> mosaics

Your dataset is an observation of the Southern region



With increasing frequency:

★ No external human interferences in the data



- **★** No ionospheric effect
- * Tropospheric effects: absorption and delay of signal
 - stronger weather dependency
 - T_{sys} dominated by atmospheric noise



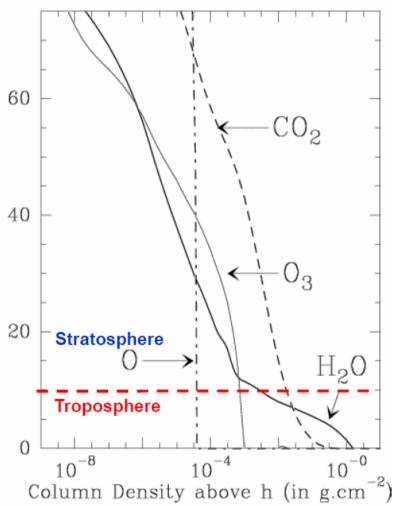


The role of the troposphere

Column density as function of altitude

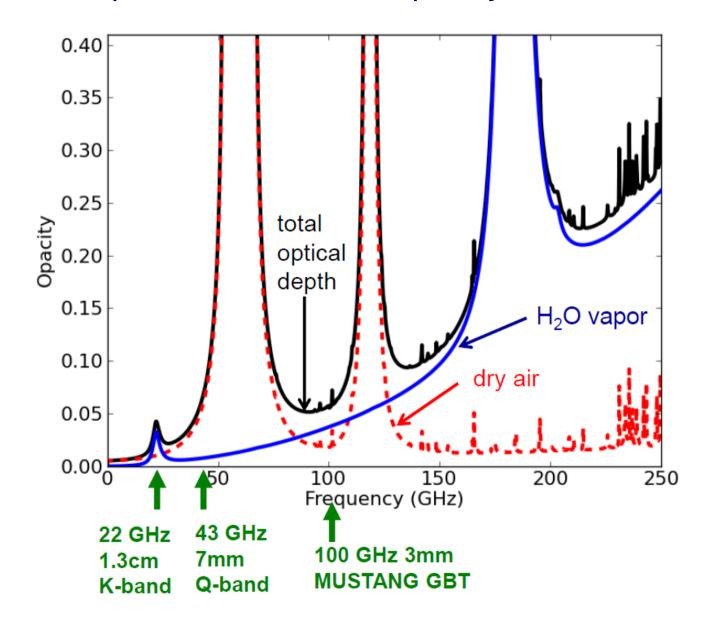
- H₂O (mostly vapor)
- "Hydrosols" (water droplets in clouds and fog)
- "Dry" constituents: O₂, O₃, CO₂, Ne, He, Ar, Kr, CH₄, N₂, H₂

clouds & convection = time variation



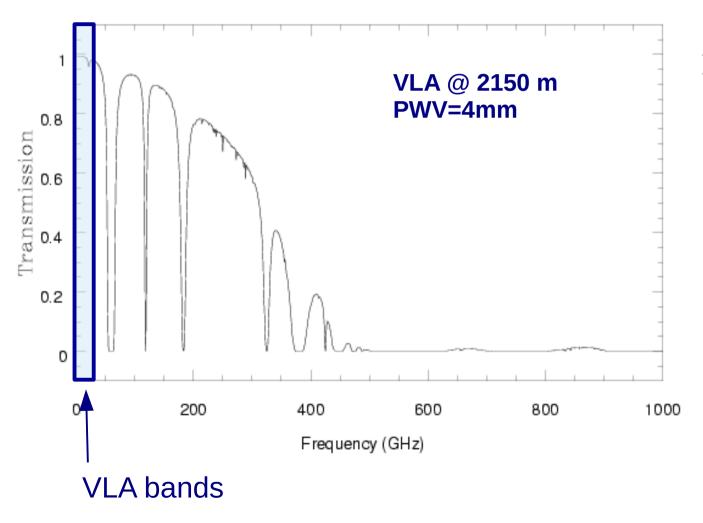


Optical depth as function of frequency





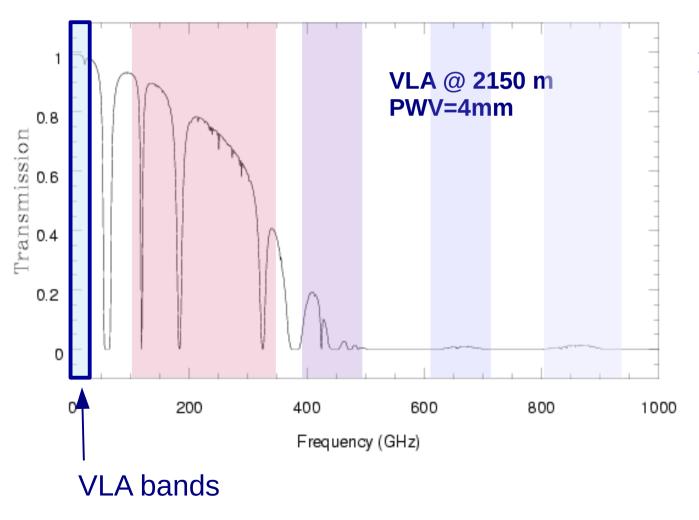
Tropospheric opacity depends on altitude



Atmospheric transmission not a problem $@ \lambda > cm$



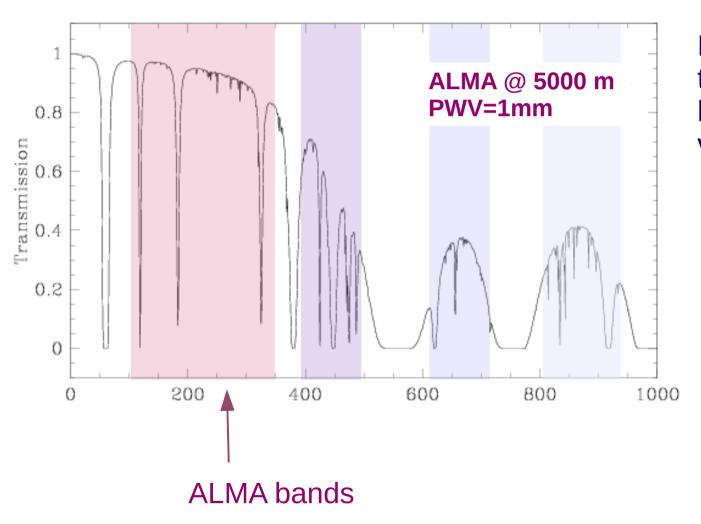
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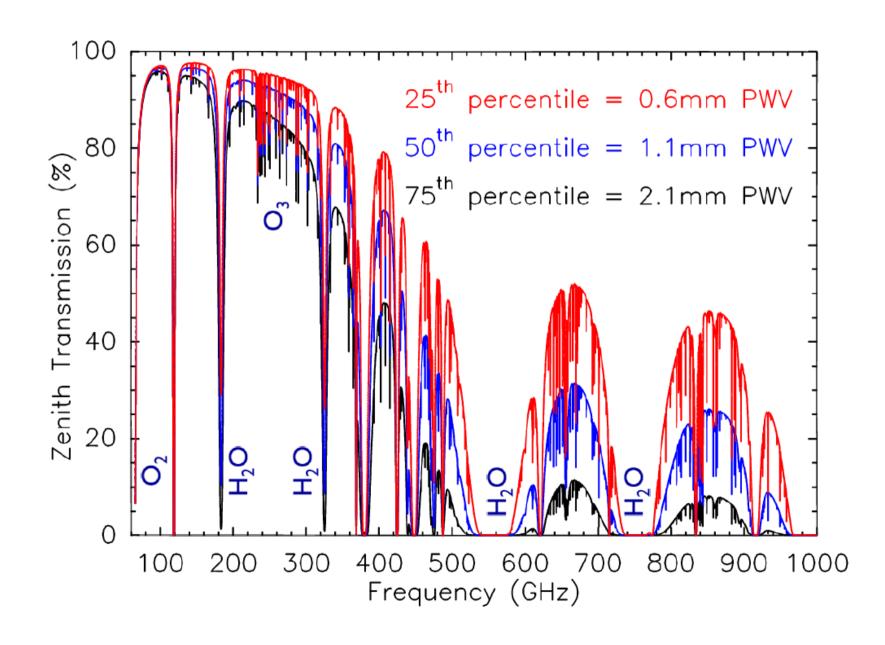
Tropospheric opacity depends on altitude



Difference due to the scale height of water vapor

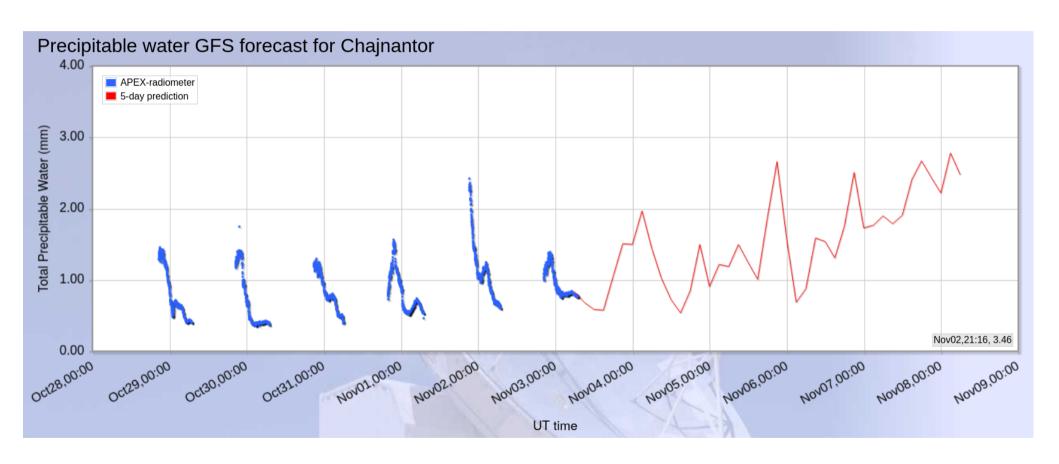


PWV= Precipitable Water Vapour



Dynamic scheduling

To efficiently use ALMA's capabilities under varying environmental conditions:



http://www.apex-telescope.org/weather/RadioMeter/index.php

Dynamic scheduling

To efficiently use ALMA's capabilities under varying environmental conditions, an observation is divided in blocks of self-consistent observations **EBs** "Execution blocks"

The project we are working on has 6 EBs:

```
uid ___A002_X1ff7b0_X1c8
uid __ A002_X207fe4_X1f7
uid __ A002_X207fe4_X4d7
uid __ A002_X215db8_X18
uid __ A002_X215db8_X1d5
uid __ A002_X215db8_X392
```

Each of them contains all the observations of the calibrators needed to properly calibrate the scientific data



Mean effect of atmosphere on Phase

Variations in precipitable water vapor (PWV) cause phase fluctuations, worse at higher frequencies, resulting in:

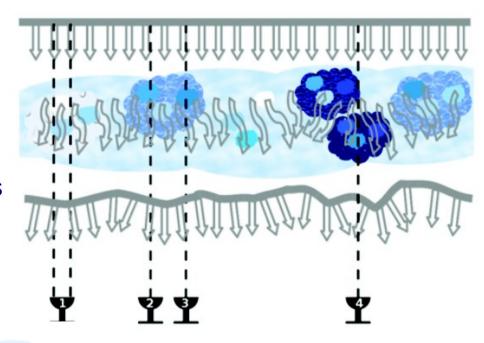
- Phase shift due to refractive index $n \neq 1$
- Low coherence (loss of sensitivity)

Patches of air with different pwv (and hence index of refraction) affect the incoming wave front differently.

Antenna 1, 2, 3 see slightly different disturbances

Sky above antenna 4 varies independently

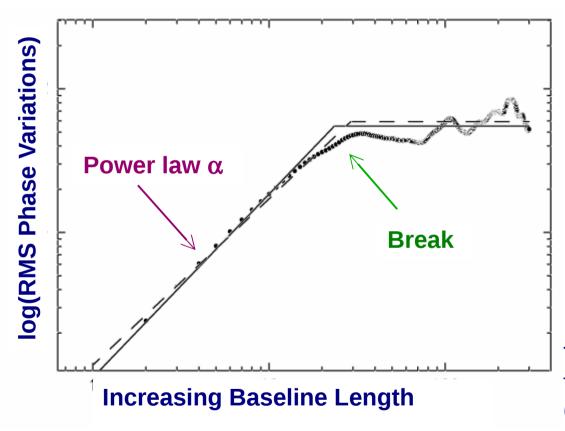
The phase change experienced by an e.m. wave can be related to pwv



$$\varphi_e \approx \frac{12.6 \,\pi}{\lambda} \cdot pwv$$



Atmospheric phase fluctuations



Phase noise

 $\varphi_{rms} = \frac{K b^{\alpha}}{\lambda}$

Kolmogorov turbulence theory

b=baseline length (km)

 α = 1/3 to 5/6 (thin or thick atmosphere)

 λ = wavelength (mm)

K constant (~100 for ALMA)

The break is typically @ baseline lengths few hundred meters to few km (scale of the turbulent layers)

Break and maximum are weather and wavelength dependent



Atmospheric phase fluctuations → decorrelation

We lose integrated flux because visibility vectors partly cancel out

$$\langle V \rangle = V_o \langle e^{i\varphi} \rangle = V_o e^{-(\varphi_{rms}^2)/2}$$

$$\phi_{rms} = 1 \text{ radian } \rightarrow \langle V \rangle = 0.60 V_0$$

In summary

Fluctuations in the line-of-sight pwv of an antenna cause phase variations of the order of ~30 deg / sec at 90 GHz, and scales linearly with frequency....

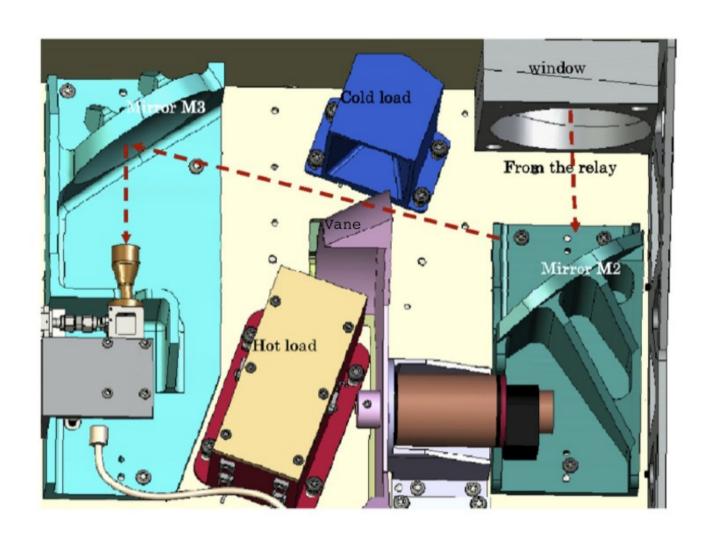
 $\varphi_e \approx \frac{12.6 \,\pi}{\lambda} \cdot pwv$

and the phase noise is worse at longer baselines...

$$\varphi_{rms} = \frac{K b^{\alpha}}{\lambda}$$



Each ALMA 12 m antenna has a water vapour radiometer

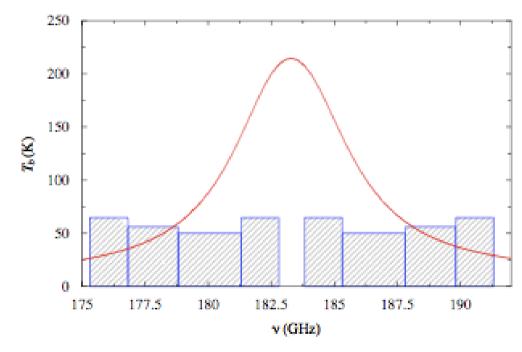




Each ALMA 12 m antenna has a water vapour radiometer

Four "channels" flanking the peak of the 183 GHz water line

Data taken every second





Each ALMA 12 m antenna has a water vapour radiometer

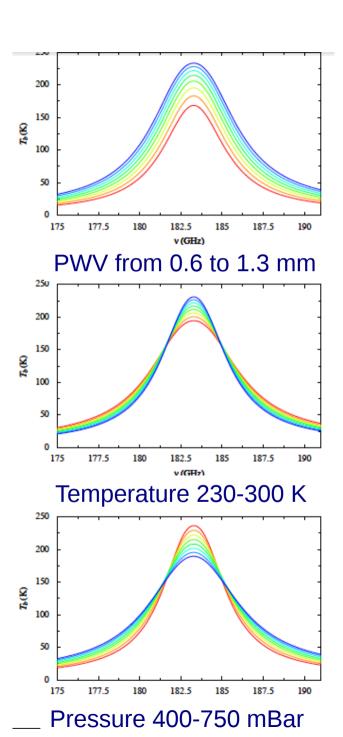
Four "channels" flanking the peak of the 183 GHz water line

Data taken every second

Convert 183 GHZ brightness to PWV (wvrgcal):

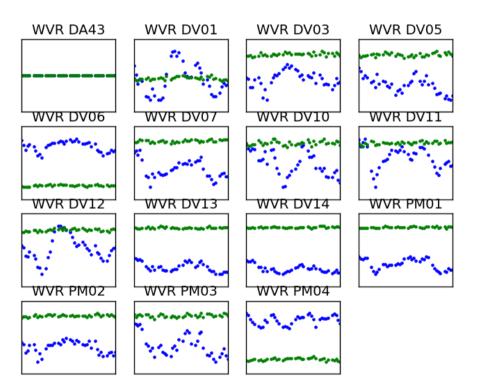
model PWV, temperature and pressure compare to the observed "spectrum" compute the correction:

$$\varphi_e \approx \frac{12.6 \,\pi}{\lambda} \cdot pwv$$



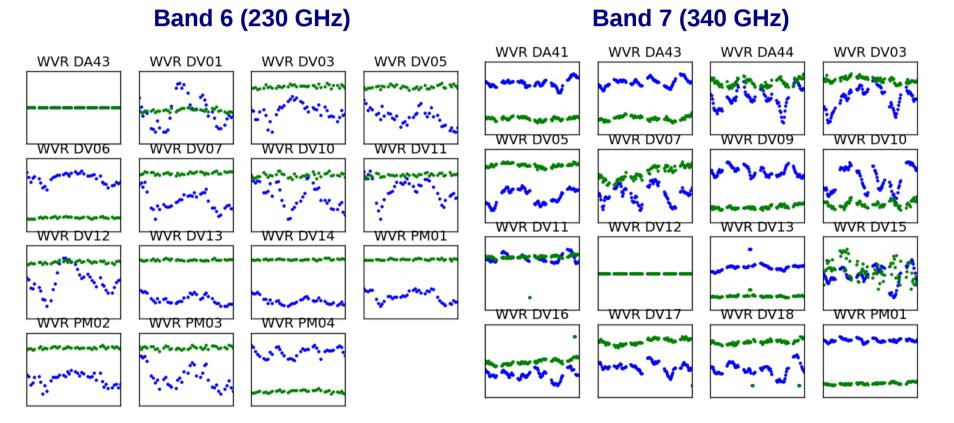


Band 6 (230 GHz)



Raw phases & WVR corrected phases





Raw phases & WVR corrected phases

e.g. to observe a 1 Jy source with a 10 m radiotelescope we have to measure $T_A \sim 0.04$ K against $T_{sys} \sim 100$ K

$$T_{sys} \sim T_{atm} (1 - e^{-\tau}) + T_{rx}$$

At lower frequencies T_{rx} is dominant

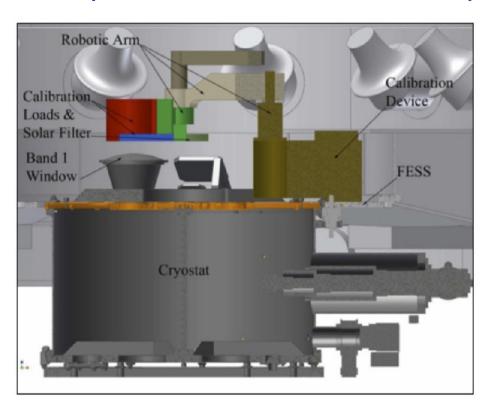


At higher frequencies (mm/submm) the noise associated with the atmosphere T_{atm} is dominant, and acts like a blackbody emitter, attenuating the astronomical signal



System noise temperature

ALMA front end are equipped with an Amplitude Calibration Device (ACD)



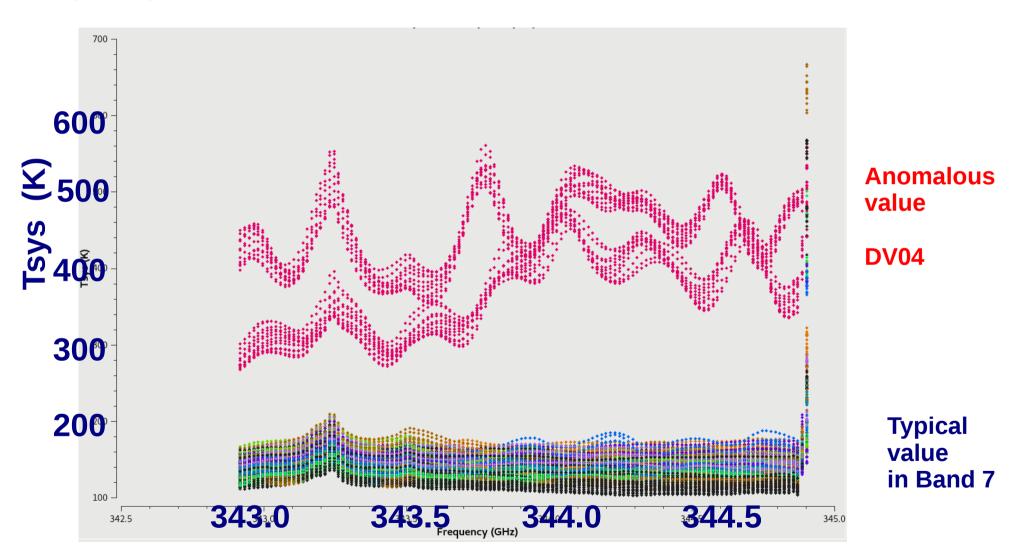


To measure $T_{sys} \text{ and } T_{rx}$ stored in tables

Every scan could have a Tsys measurement, but <400 GHz relatively constant ~10min.
Tsys spectra are applied off-line to the correlated data.

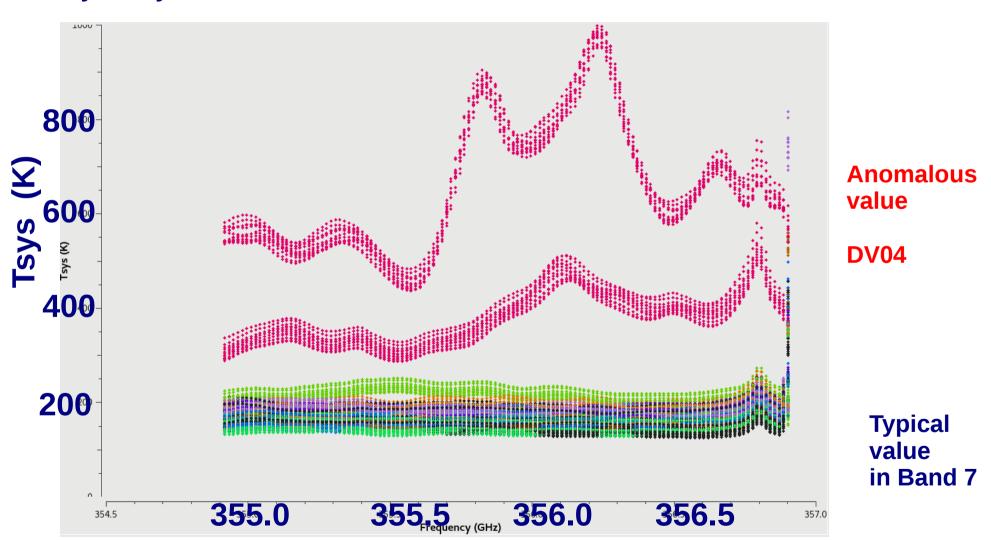
Assuming correlated data in units of % correlation multiplication by Tsys will change the unit to Kelvin

Tsys in your dataset: in color different antennas



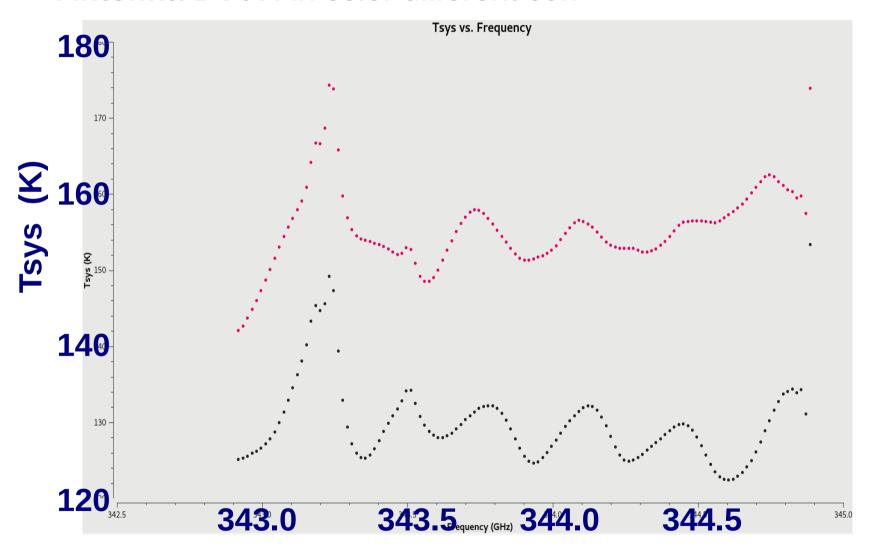
Frequency (GHz)

Tsys in your dataset: in color different antennas



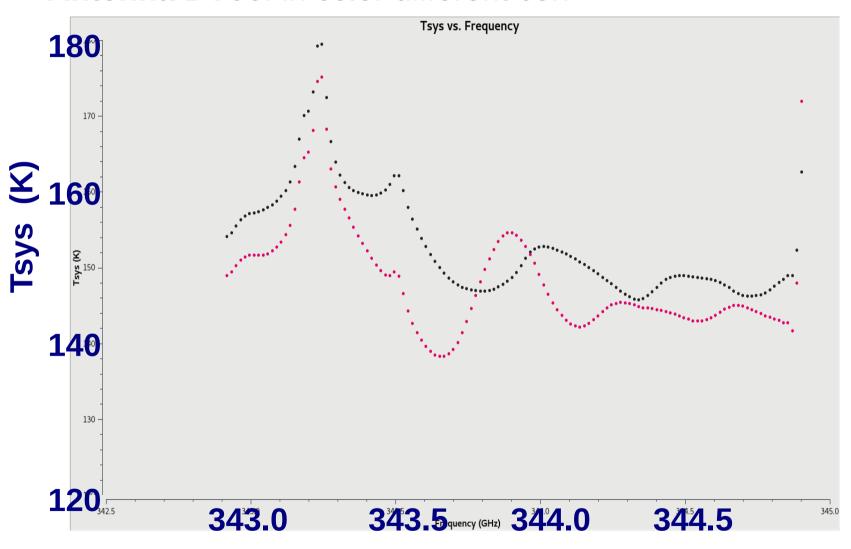
Frequency (GHz)

Antenna DV07: in color different corr



Frequency (GHz)

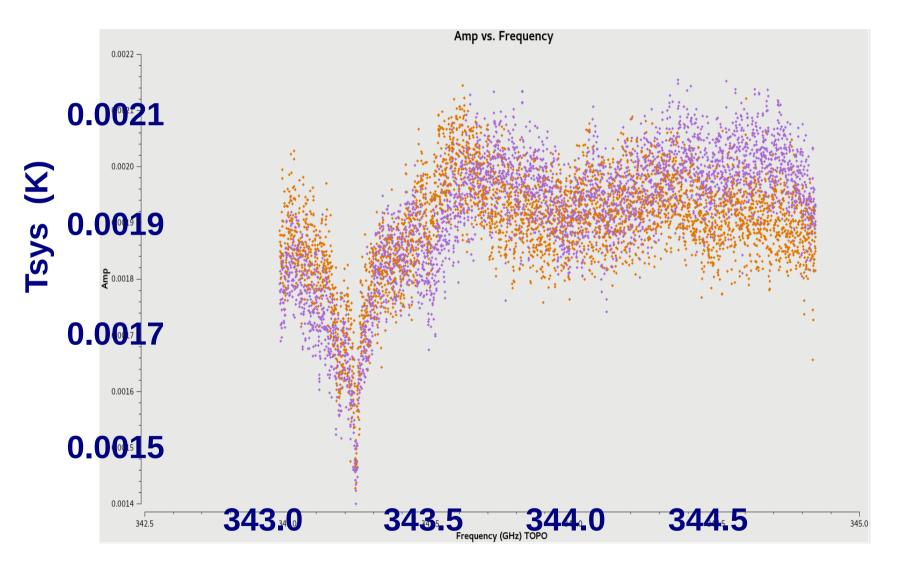
Antenna DV06: in color different corr



Typical value in Band 7

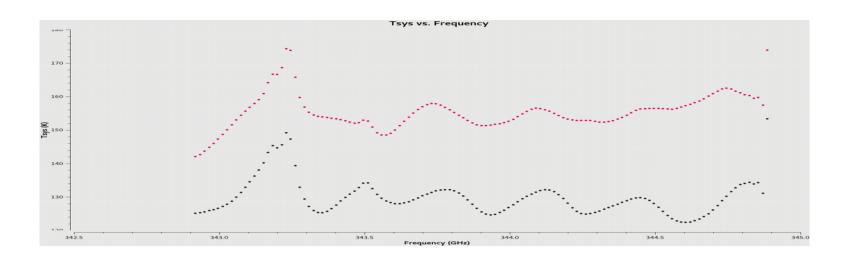
Frequency (GHz)

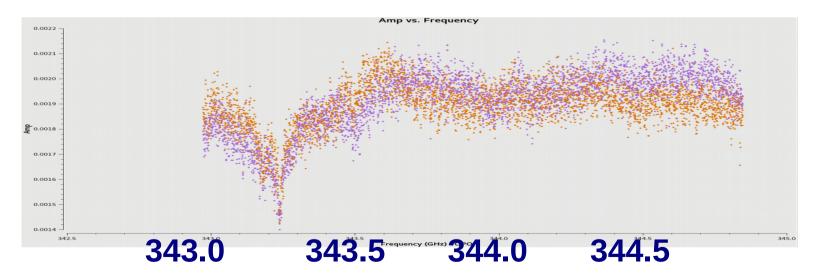
Baseline DV06&DV07: in color different corr



Frequency (GHz)

Baseline DV06&DV07: in color different corr



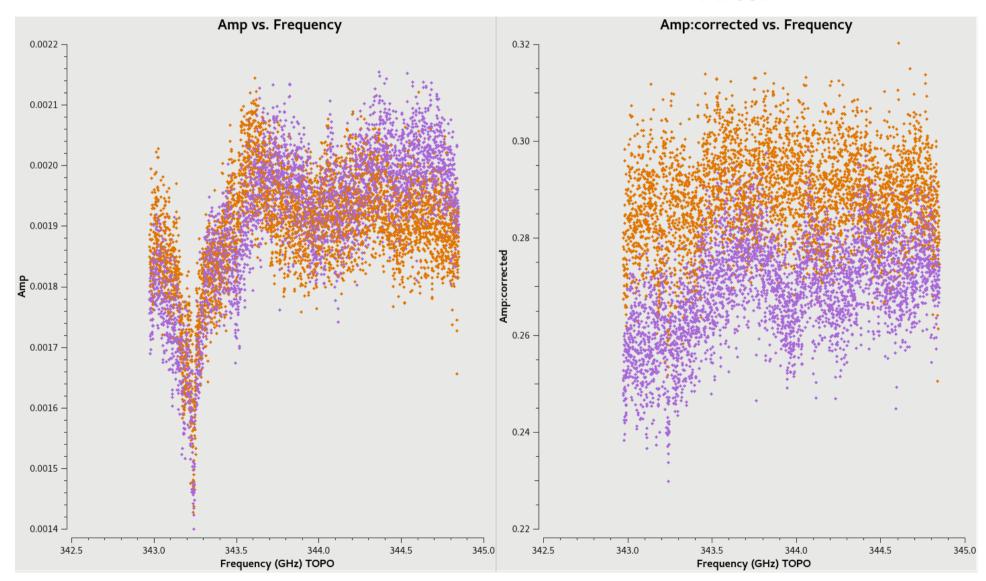


Frequency (GHz)

Tsys calibration DV06 & DV07 Before

The attenuation Is corrected

After



Tsys calibration DV06 & DV07 Before



The data are now Temperature in K

After

